

"4.1.BMP"

CALCULATE THE GRAVITY ANOMALY DUE TO THE LOAD OF HAWAII AND ESTIMATE THE ELASTIC THICKNESS OF THE OCEANIC LITHOSPHERE, T_e.

This Mathcad file calculates the gravity associated with the flexure of the oceanic crust by the load of the Hawaiian-Emperor seamount chain and compares it to the observed free-air gravity anomaly in the region of Oahu. The elastic thickness and load and infill density that best explains the observed gravity anomaly can be determined from the comparison.

<u>Define input parameters</u>

Define elastic thickness, T_e , average gravity, g, Poisson's ratio, v, Youngs Modulus, E, and the Universal Gravitational constant, G.

Te := 25km $g_{\rm a}$:= 9.81m·sec⁻² v := 0.25 E := 10¹¹Pa $G_{\rm c}$:= 6.67·10⁻¹¹·newton·m²·kg⁻²

Define npts, the number of sample points along the gravity and topography profile. npts must be a

power of 2 (for the Fourier Transform).

npts := 128

Define density of the load (i.e. the seafloor topography), water and mantle

 $\rho_{\text{crust}} := 2800 \text{kg·m}^{-3}$ $\rho_{\text{water}} := 1030 \text{kg·m}^{-3}$ $\rho_{\text{mantle}} := 3330 \text{kg·m}^{-3}$

Define normal (ie unflexed) thickness of oceanic crust

ocean_thick := 6km

$D := \frac{E \cdot (T e)^3}{12 \cdot \left(1 - v^2\right)}$

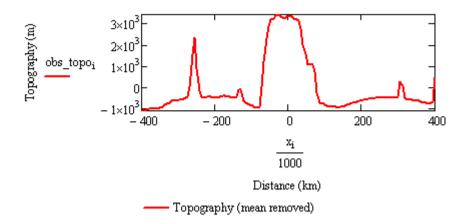
Read in the observed topography and free-air gravity anomaly data

$$\begin{split} \mathbf{M} &\coloneqq \text{READPRN}(\text{"haw_topo.pm"}) \\ \underbrace{\mathbf{N}_{i:}}_{i:ountt} &\coloneqq \text{READPRN}(\text{"haw_grv.pm"}) \\ \text{icountt} &\coloneqq \text{rows}(\mathbf{M}) - 1 \\ \text{profile} &\coloneqq \left(\mathbf{M}_{icountt,0} - \mathbf{M}_{0,0} \right) \text{km} \\ j &\coloneqq 0... \text{rows}(\mathbf{M}) - 1 \\ a_{j} &\coloneqq \mathbf{M}_{j,0} \text{ km} \\ b_{j} &\coloneqq \mathbf{M}_{j,1} \text{ m} \\ j &\coloneqq 0... \text{rows}(\mathbf{N}) - 1 \\ \underbrace{\mathcal{C}_{j}}_{j} &\coloneqq \mathbf{N}_{j,0} \text{ km} \\ d_{j} &\coloneqq \frac{\mathbf{N}_{j,1}}{10^{5}} \text{m} \cdot \text{sec}^{-2} \end{split}$$

Interpolate the observed data and take out a mean (for the Fourier transform)

$$\begin{split} dx &:= \frac{\text{profile}}{\text{npts} - 1} \\ XKINT &:= \frac{2 \cdot \pi}{\text{npts} \cdot dx} \\ XINT &:= \frac{\text{profile}}{\text{npts} - 1} \\ i &:= 0 \dots \text{npts} - 1 \\ i &:= 0 \dots \text{npts} - 1 \\ x_i &:= a_0 + (i \cdot XINT) \\ \text{obs_topo}_i &:= \text{linterp}(a, b, x_i) \\ \text{obs_grv}_i &:= \text{linterp}(c, d, x_i) \\ \text{mean_depth} &:= -\text{mean}(\text{obs_topo}) \\ i &:= 0 \dots \text{npts} - 1 \\ \text{obs_topo}_i &:= \text{obs_topo}_i + \text{mean_depth} \\ \text{mean_gravity} &:= \text{mean}(\text{obs_grv}) \\ i &:= 0 \dots \text{npts} - 1 \\ \text{obs_grv}_i &:= \text{obs_grv}_i - \text{mean_gravity} \\ \text{mean_depth} &= 3.825 \times 10^3 \cdot \text{m} \end{split}$$

Display observed topography data



<u>Calculate the flexural response function, Phi, and the gravitational admittance based on a</u> <u>flexure model, Zflexure</u>

$$c := fft(obs_top)$$

$$k := 0, 1 \dots \frac{npts}{2}$$

$$Phi_k := \left[\frac{D \cdot (k \cdot XKINT)^4}{g \cdot (\rho_mantle - \rho_crust)} + 1\right]^{-1}$$

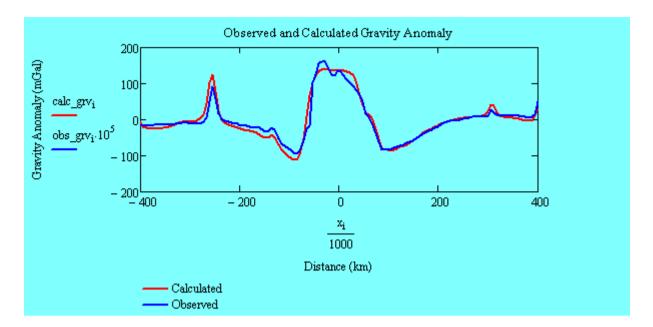
(See Eq. 5.18)

$$Z flexure_{k} \coloneqq 2 \cdot \pi \cdot G \cdot 10^{5} \cdot (\rho_crust - \rho_water) \cdot e^{-(k \cdot XKINT \cdot mean_depth)} \cdot \left[1 - \left(Phi_{k} \cdot e^{-XKINT \cdot ocean_thick}\right)\right] = 0$$

(See Eq. 5.19)

 $factor_k := Zflexure_k \cdot c_k$ calc_grv := iffl(factor)

Display results



Note: the "best fit" value of Te is one that explains both the amplitude and wavelength of the observed gravity anomaly.

Calculate correlation coefficient, r, between observed and calculated gravity anomalies

 $r := corr(obs_grv, calc_grv)$ r = 0.972

Calculate variance and RMS error between observed and calculated gravity anomalies

diff := obs_grv - $\frac{calc_grv}{10^5}$

Convert to m.sec⁻²

variance := var(diff) rms := $\sqrt{variance}$ rms = 1.491 × 10⁻⁴ ·m·s⁻²

Multiply by 10⁵ to get mGal

This gives the RMS difference (and correlation coefficient) between observed and calculated gravity anomalies based on the values of elastic thickness and densities assumed above.